# Atmospheric and Surface Water Isotopes: Processes and Applications

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## Why Water in the Atmosphere?

- Precipitation is the source of freshwater in hydro- and eco-systems
- Most of the isotopic variation we see in freshwater systems is derived from atmospheric processes



#### **The Global Meteoric Water Line**



Clark and Fritz, 1997; GMWL defined by Craig, 1961

# **Rayleigh Distillation**

- A processes by which fractionation (of isotopes, elements, molecules, elephants) leads to a change in 2component reactant mixture
- Is it Rayleigh Distillation?
  - Open system
    - No addition of material
    - Rapidly and continuous removal of product
  - Product is fractionated relative to the reactant
- The Rayleigh equation describes the composition of the reactant pool as a function of reaction progress:

#### $\boldsymbol{R} = \boldsymbol{R}_{\boldsymbol{0}} \boldsymbol{f}^{(\alpha - 1)}$

**R** and **R**<sub>0</sub> are current and initial isotope ratios **f** is fraction of reactant remaining

#### **Rayleigh Distillation**



#### **Rayleigh Distillation**



#### Rayleigh Distillation and Precipitation

 Precipitation formed from condensation of cloud vapor •Equilibrium process (free atmosphere RH ≈ 100%) •Equilibrium fractionation But one element of complexity...



#### Rayleigh Distillation and Precipitation

*!!!This does not cause the 'temperature effect'!!!* 

As condensation proceeds, the temperature of the remaining cloud decreases.

> Thus,  $\alpha$  increases, resulting in a greater difference between cloud and rain H<sub>2</sub>O



- 🔹 at 20° C
  - ε<sup>2</sup>Η = 74‰
  - $\epsilon^{18}O = 9.2\%$
  - $\epsilon^{2}H / \epsilon^{18}O = 8.0$
- 🔹 at 80° C
  - $\bullet \epsilon^2 H = 38\%$
  - $\epsilon^{18}O = 4.5\%$
  - $\epsilon^{2}H / \epsilon^{18}O = 8.4$

Equilibrium enrichment factors for H isotopes are ~8 x those for O isotopes

#### **The Global Meteoric Water Line**



Clark and Fritz, 1997; GMWL defined by Craig, 1961

#### Phase Change Reaction: Craig-Gordon

- Open air
  - Well-mixed
  - Large
- Transition zone (TZ)
  - Turbulently mixed
  - Decreasing humidity upwards
- Boundary layer (BL)
  - Thin, well-mixed layer
  - 100% RH
- Liquid
  - Large (ocean) or small (droplet) body of water
  - Mixed or stratified



#### **Kinetic Effect**

#### • Δε = (1-h)n $\theta C_D$

- h = relative humidity
- n = relative strength of kinetic vs. turbulent transport
- $\theta$  = perturbation of boundary layer humidity
- C<sub>D</sub> = ratio of effective diffusion coefficients for isotopologues
  - H = 25.1‰
  - O = 28.5‰

#### Let's discuss...

# Equilibrium and kinetic enrichment factors



Gat, 1996

# Equilibrium and kinetic enrichment factors

- The net ratio of <sup>2</sup>H and <sup>18</sup>O isotope effects is a blend of the Equilibrium and Kinetic ratios, typically between 3 and 8
- The coupled <sup>2</sup>H/<sup>18</sup>O system gives us a "proxy" for kinetic fractionation...

deuterium excess





#### **Deuterium excess**

#### $d = \delta^2 H - 8 \times \delta^{18} O$



Crystal Tulley-Cordova, unpub.

#### **GMWL Intercept**

- Global precipitation
  - δ<sup>2</sup>H = -22‰
  - $\delta^{18}O = -4\%$
- What is the isotopic composition of global evaporation?
- So global evaporation has a d value of ~ +10‰
  - Implies conditions of evaporation
    - n = 0.5
    - $\theta = 0.5$

#### **The Global Meteoric Water Line**



Clark and Fritz, 1997; GMWL defined by Craig, 1961

The slope of the GMWL is 8

Kinetic fractionation gives a *d* value for the (dominant) oceanic evaporation flux of +10‰, setting GMWL intercept

#### **Precipitation Isotope Patterns and Rayleigh Distillation**

 Rayleigh distillation causes precipitation isotope ratios to get lower as air gets drier

What does this imply about the spatial distribution of precipitation isotope ratios?

# Patterns of Climate $\rightarrow$ Patterns of Water Isotope Ratios



#### A Salt Lake City Example



#### **Temporal Variation - Meterological**



Coplen et al., 2008, GRL

## Synoptic-scale systems (Sandy)

October 22-31, 2012
\* ET transition October 29
\* Landfall (NJ) early on October 30
Maximum intensity category 3, category 1 at landfall
2nd costliest Hurricane in US history





## **Crowdsourced Sampling Network**



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#### **Sample Suite**



Good et al., PLoS ONE, 2014

#### **Spatiotemporal Isotope Patterns**



Distinct regional isotopic patterns as storm moves northwest

## **Rayleigh Rainout Model**



## Spatial Evolution of Storm Water Cycling



- Superstorm Sandy up to 80% rainout-`efficient'
- Over 2 days disconnects from Atlantic, adds continental moisture

Good et al., 2014, Sandy edited volume

## **Pacific North American Pattern**



http://www.emc.ncep.noaa.gov/gmb/ssaha/

#### PNA Pattern and Precipitation δ<sup>18</sup>O



Liu et al., 2013, Clim. Dynamics

#### **A PNA Isotope Index**



*Liu et al., Earth and Planetary Science Letters, 2011* 

**PNA** index

#### **Paleo-PNA**



Liu et al., Nat. Comm., 2014

#### **Multiple Sources of ET**



#### Water Isotope Tracers

- Stable H and O isotopes inherent tracers with welldocumented, well-understood behavior in the water cycle
- Transpiration does not fractionate

•  $\delta_{vapor} = \delta_{liquid} = \delta_{precipitation}$ 



## **2** Approaches to ET Separation

- Measure vapor
  - + Directly assess ET flux
  - Difficult at large scales



Griffis et al., 2010, *B-LM*
# **2** Approaches to ET Separation

#### Measure liquid

- + Can integrate over large areas
- May not `see' all processes



# **Ecohydrologic Connectivity**

 Water in soils and plants may not be same water that enters aquifers and rivers



Brooks et al., 2009, Nat. Geoscience

#### **Can we Combine Approaches?**

- $\ensuremath{^{\scriptsize \ensuremath{\bullet}}}$  Need global estimates of  $\delta$  for river water and ET
- Where can we get them?
  - Ocean and atmosphere mass-balance

$$\begin{split} dH_2O_0/dt &= 0 = P_0 + Q - E_0 \\ d\delta_0/dt &= 0 = P_0^*\delta_{P_0} + Q^*\delta_Q - E_0^*\delta_{E_0} \\ dH_2O_a/dt &= 0 = E_0 + ET - P_0 - P_c \\ d\delta_a/dt &= 0 = E_0^*\delta_{E_0} + ET^*\delta_{E_T} - P_0^*\delta_{P_0} - P_c^*\delta_{P_C} \end{split}$$

#### Constraints

Good data exist on bulk water fluxes (P, E, ET, Q)
 Precipitation isotope composition from isoscapes



### **The Evaporation Isotope Problem**

 Excellent model for estimating δ value of E<sub>o</sub>, but you need to know surface layer δ<sub>a</sub>



*Figure 3* The Craig-Gordon evaporation model.

Gat, 1996, AREPS

#### New Satellite-Based $\delta_a$ Isoscape



Good et al., 2015, GRL

# New Estimates of Global $\delta_{Q'}$ , $\delta_{Eo}$



Good et al., 2015, GRL

# **Global Analysis**

What combinations of T, soil E, open-water E, and connectivity give the estimated global values

Good, Noone and Bowen, 2015, Science



#### **Kernel Density Examples**



Good, Noone and Bowen, 2015, Science

# **Global Result**

- Globally, transpiration majority of ET but not 90%
- Low connectivity of soil and surface waters suggests dominance of preferential flow
- This, + the fact that >60% of continental evaporation occurs in soils, compromises the lakebased ET separation method

Good, Noone and Bowen, 2015, Science



#### **Catchment water balance**



 $P = R + E + T + \Delta S$ 

# Storm Hydrograph

- Accurate rainfall-runoff models are one of the fundamental goals of catchment hydrology
  - Flood control
  - Water management
  - Water quality
- Summarized in terms of storm hydrograph



# **Rainfall routing**

How is storm precipitation delivered to streams?

- Overland flow
- Interflow
- Groundwater
  recharge/baseflow
- What are the transit times associated with these flowpaths?



#### www.hydro.washington.edu

# Hydrograph separation

 Given any conservative tracer that is present in different abundance in *pre-event* and *event* water, the fraction of pre-event water in storm flow at any time is

$$f_{pe} = \frac{C_s - C_e}{C_{pe} - C_e}$$

- Major assumption: C<sub>e</sub> and C<sub>pe</sub> are constant and can be accurately characterized
- Isotope hydrograph separation takes advantage of temporal variation in precipitation δ values



# "Old water paradox"

		Catchment		Percentage pre-event water	
Study	Location	area (ha)	Tracer	peak	volume
Jordan (1994)	Switzerland	3.6	<sup>18</sup> O		45, 75
Waddington et al. (1993)	Ontario	160	<sup>18</sup> O	87,93	
McDonnell et al. (1991)	New Zealand	3.8	D	92-100	
O'Gunkoya & Jenkins (1991)*	United Kingdom	1000	D		54-90
McDonnell et al. (1990)	New Zealand	310	D		21-33
Nolan and Hill (1990)	California	1060	D		57
Bonell et al. (1990)	New Zealand	218	D		59
		310	D	38	38 to >97
Blowes and Gillham (1988)	Ontario	0.75	<sup>18</sup> O	9, 45	22, 50
Turner et al. (1987)	W. Australia	. 82	<sup>18</sup> O, D		69-95
Herrman et al. (1987)*	Germany	76	<sup>18</sup> O		84
Rodhe (1987)	Sweden	3	0 <sup>81</sup>		81, 87
		4	<sup>18</sup> O		81, 96
	· · · ·	17	18 <sub>0</sub>	87	
		50	<sup>18</sup> O		85-99

# "Old water paradox"

- Even in small, steep catchments, most storm flow is "old" water
- Mean residence time of water in catchments is much longer than implied by simple interpretation of storm hydrograph, rainfall-runoff models
- How does storm discharge actually work?
  - If water is not transferred directly, how does addition of precipitation rapidly force release of pre-event water?
- Hypotheses
  - Pressure waves
  - Capillary fringe
  - Macropore flow

### **Baseflow transit time distributions**

 Given measured input and output time series, optimize the function f(t) describing the distribution of transit times within a catchment



### **Baseflow transit time distributions**





### **Frontiers: Managed Landscapes**

- Relatively few studies have investigated transit times and runoff routing in large catchments and humandominated systems
- LOTS of great questions relating to land-use effects on runoff generation



# Hydrograph separation - suburban







Pre-event water
 <33% of storm</li>
 runoff



#### **Urban system – Red Butte Creek**

 Rapid and substantial addition of storm water in lower developed reach of catchment (UU campus)



# Agricultural systems – artifical drainage

- Subsurface drain network increases agricultural land quality
- Fundamentally alters hydrological flow
- What is the impact on timing, magnitude of nutrient discharge from these lands?





# **Hoagland watershed**

- Paired catchment study
- High and Low drainage density catchments
- Sampling for water quality, isotopes through Nov. storm
- Student participation (EAS591 Isotope Hydrology)





# Storm hydrograph

 Different timing of discharge for the highand low-drainage catchments





# Hydrograph separation

- Partitioned storm flow into event and pre-event water components using O isotopes, Mg<sup>2+</sup>
- Faster routing of storm water to stream in high-drainage catchment
- Change in storm water routing through event



# **Nutrient discharge**

 Timing and magnitude of peak nitrate discharge different in the two catchments



# Flush/bypass model

- Late season, drains largely inactive
- Drains divert storm water, slowing water table recharge
- Drains activation and discharge of high-N groundwater lags in high-drainage catchment
- Diversion of infiltration changes 'flashiness' of contaminant export



# River water isotopes – continental scale



Kendall & Coplen, 2001

# River water isotopes: rainfall-runoff model

 Can we reproduce isotope differences between surface water and precipitation using process models?







Bowen et al., 2011, JGR



#### **Runoff model**



Predicted river  $\delta^2 H$  values superimposed on precipitation  $\delta^2 H$  values

#### **Runoff model residuals - why?**



Data: Kendall and Coplen, 2001

# Sensitivity testing ET patterns

#### Impact of accounting for different % loss of precipitation to ET by month



Bowen et al., 2011, JGR
# Interrogating water source

- Evaporation
- Winter-biased runoff
- Imported water
- Provides basis for quantifying (and monitoring) seasonal or elevation bias of runoff generation
- Identifies non-local water sources



# Managed waters

- Where does your water come from?
  - The faucet
  - 🕈 Yes, but...
- With spatial information isotopes can
  - Establish connectivity
  - Document what happens between the source and sink



## **US** tap water isotope ratios

#### Range

- + 152‰ < δ<sup>2</sup>H < +11‰</li>
- -19.4‰ <  $\delta^{18}$ O < +4.2‰
- Cluster near the GMWL
- Average d-excess value of 5 not significantly different from 10



Bowen et al., 2007, WRR

#### **Isotope ratios of US tap waters**





#### Tap water vs. local precipitation

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### Tap water d values compared to precipitation



# Catchment areas: source/sink connectivity

- Characterize source region for water supply
- Map supply footprint of source
- Water sinks 1000+ km removed from mean sources



# **Scaling Down: Urban Structure**

- Many urban centers have decentralized water management
- Do importation rules scale down from states to cities?
- Inference: how are resources being managed, what are controls
- Prediction: neighborhoodscale predictive isoscapes?



# Salt Lake City Tap Water Isotopes



#### **Multiple Modes of Variation**



# Spatiotemporal Water Isotope Patterns

